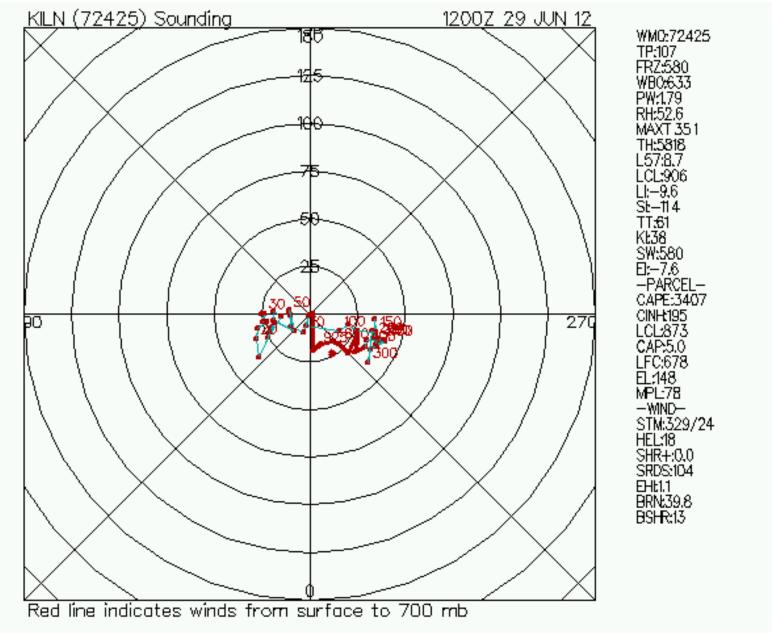
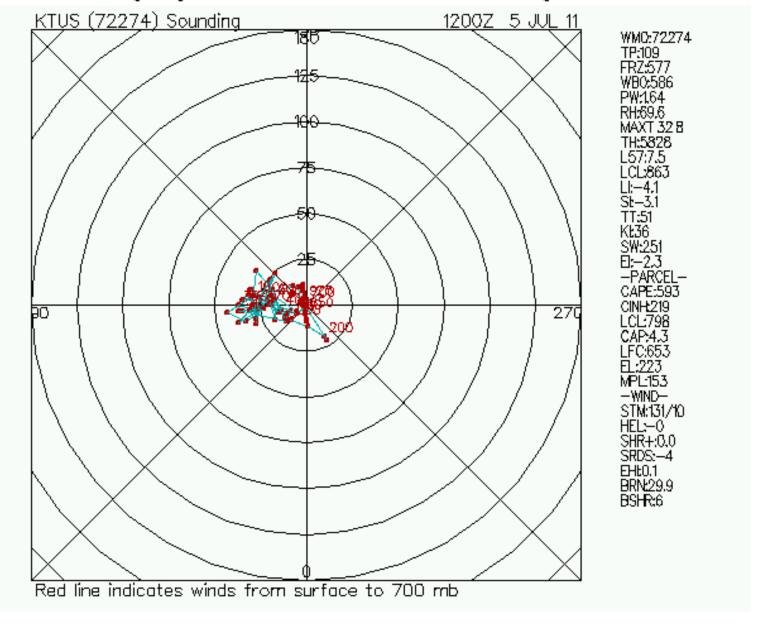
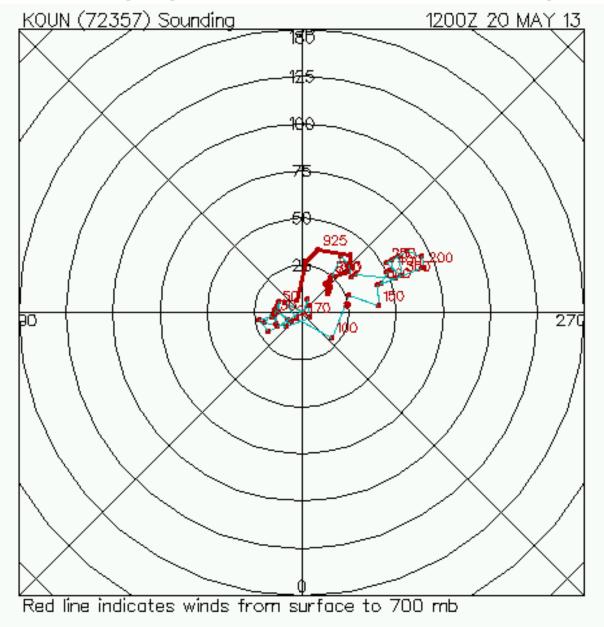
▼ Plymouth State Weather Center **▼**



▼ Plymouth State Weather Center **▼**



▼ Plymouth State Weather Center **▼**



WM0:72357 TP:210 FRZ:620 WB0:687 PW:100 RH:43.5 MAXT 34.6 MAXT 34.5 TH:5722 L57:8.5 LCL:940 LI:-9.1 St-2.5 TT:55 KI:24 SW:364 EI:-19 -PARCEL-CAPE:2146 CAPEZIA CINH274 LCL:874 CAP:6.7 LFC:699 EL:199 MPL:119 -WND-STM:260/25 HEL:327 SHR+:0.0 SRDS:68 EHt4.9 BRNt39.6 BSHR:10

1 Hodograph

Hodographs represent the vertical profile of wind direction and magnitude captured from soundings. Each point along the hodograph corresponds to wind at a given level; the distance from the origin shows the magnitude and the angle to the origin gives the wind direction. Mean winds and shear characteristics for a given layer can be estimated for a given layer by visual inspection of the hodograph. The length of the hodograph (or part of the hodograph if considering a certain layer) corresponds to the total speed shear, curvature of the hodograph corresponds to vertical changes in the direction of the shear vector $\mathbf{S} = \frac{d\mathbf{v}}{dz}$.

2 Shear vector and horizontal vorticity

The shear vector **S** at a given level can be approximated from sounding data by finite difference, $\mathbf{S} \approx \frac{\mathbf{v}_2 - \mathbf{v}_1}{z_2 - z_1}$, has units of \mathbf{s}^{-1} , and in the absence of significant horizontal gradients of vertical motion, can be rotated -90° to obtain horizontal vorticity, $\omega_h = \hat{\mathbf{k}} \times \mathbf{S}$.

3 Streamwise and crosswise horizontal vorticity

Sounding data and the information contained in the hodograph are used to characterize the environment in which convective storms will develop. Storm-relative winds are obtained by estimating a storm velocity c and subtracting from observed winds, $\mathbf{v}_{SR} = \mathbf{v} - \mathbf{c}$. I estimated c as the mean winds in the 850-600 mb layer. The storm-relative wind at a given layer can then be used to decompose the horizontal vorticity into streamwise vorticity, $\omega_s = \frac{\mathbf{v}_{SR}}{\|\mathbf{v}_{SR}\|} \cdot \omega_h$, and crosswise vorticity, $\omega_c = (\hat{\mathbf{k}} \times \frac{\mathbf{v}_{SR}}{\|\mathbf{v}_{SR}\|}) \cdot \omega_h$ which is useful to understanding how parcels will behave when caught in storm updrafts.